USING MGS TES DATA TO UNDERSTAND WATER CYCLING IN MARS' NORTH POLAR REGION.

L. K. Tamppari¹, A. S. Hale², D. S. Bass³ and M. D. Smith⁴, ¹NASA Jet Propulsion Laboratory (MS 301-422, 4800 Oak Grove Drive, Pasadena, CA 91109 leslie.k.tamppari@jpl.nasa.gov), ²NASA Jet Propulsion Laboratory (MS 264-235, 4800 Oak Grove Drive, Pasadena, CA 91109 amy.s.hale@jpl.nasa.gov), ³NASA Jet Propulsion Laboratory (MS T1722, 4800 Oak Grove Drive, Pasadena, CA 91109 deborah.s.bass@jpl.nasa.gov) ⁴NASA Goddard Space Flight Center (Michael.D.Smith.1@gsfc.nasa.gov).

Introduction: The Martian water cycle is one of the three annual cycles on Mars, dust and CO2 being the other two. Despite the fact that detailed spacecraft data, including global and annual coverage in a variety of wavelengths, have been taken of Mars spanning more than 25 years, there are many outstanding questions regarding the water cycle.

There is very little exposed water on Mars today, in either the atmosphere or on the surface [1] although there is geological evidence of catastrophic flooding and continuously running water in past epochs in Mars' history [2] as well as recent (within about 10,000 years ago) evidence for running water in the form of gullies [3]. While there is little water in the atmosphere, water-ice clouds do form and produce seasonal clouds caused by general circulation and by storms [5-8]. These clouds may in turn be controlling the cycling of the water within the general circulation [e. g., 6].

The north polar cap region is of special interest as the residual cap is the main known reservoir of water on the planet today. The south polar residual cap may contain water, but presents a CO2 ice covering, even during southern summer. This hemispheric dichotomy is unexplained and is especially puzzling due to the fact that the Martian southern summer is much warmer (due to Mars' eccentricity) than the northern summer. Recently, water has been found in the top meter of the surface in both the northern and southern high latitude regions [e.g., 8-9] indicating and even greater amount of water on Mars than previously known.

Background: In order to understand the polar dichotomy, our approach is to examine the water transport and cycling issues within the north polar region and in/out of the region on seasonal and annual timescales. Viking Mars Atmospheric Water Detector (MAWD) data showed that water vapor increased as the northern summer season progressed and temperatures increased, and that vapor appeared to be transported southward [10]. However, there has been uncertainty about the amount of water cycling in and out of the north polar region, as evidenced by residual polar cap visible brightness changes between one Martian year (Mariner 9 data) and a subsequent year (Viking data). These changes were originally thought to be interannual variations in the amount of frost sublimed based on global dust storm activity [10-12]. However,

Viking thermal and imaging data were re-examined and it was found that 14-35 pr mm of water -ice appeared to be deposited on the cap later in the summer season [14], indicating that some water may be retained and redistributed within the polar cap region. This late summer deposition could be due to adsorption directly onto the cap surface or due to snowfall. The possibility that some of the water is seasonally sequestered in water-ice clouds and may allow later precipitation had not been previously considered. We address these issues by examining water vapor and water-ice clouds in the north polar region of Mars during the north spring and summer period.

Method:

Water-ice clouds. Water-ice clouds, in the north polar region, have previously been tentatively identified in the Viking era using the Infrared Thermal Mapper (IRTM) data [14] and in the Mars Global Surveyor (MGS) era using the Thermal Emission Spectrometer (TES) data [15] and M. Smith, pers. comm., 2001). The Viking data provides only nadir pointed data, which necessitates separating the surface contribution from the atmospheric contribution. The technique used relies on the water-ice absorption feature at 11microns. Specifically, it uses the 11- and 20-micron channels of the IRTM instrument along with a surface model to accurately remove the surface contribution [4]. For the Viking time frame, we are restricted to ice free surfaces and therefore do not see the entire north polar cap region throughout the spring and summer season. With the MGS TES data, there are both limband nadir-pointed data. The nadir-pointed data retrievals may be accomplished for water ice clouds only over surfaces above 220 K (M. Smith, pers. comm.), but limb data do not have this limitation. Therefore, water-ice clouds can be identified throughout the Martian season, even over the seasonal and residual polar cap.

Water Vapor. Recent analysis with MGS TES data has shown evidence for water vapor "pulses" as the seasonal north polar cap sublimes [15]. This could be linked to the previous late-summer season deposition, discussed above. Additionally, there may be some differences in the details of the water vapor as a function of latitude and season between the Viking era and the current era. Any such differences would help

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identify and characterize the degree of interannual variability in the water vapor.

Dust. Dust in the north polar region can play an important role in the water cycling story in a couple of ways. One, by acting as cloud condensation nuclei, dust may allow water-ice clouds to form and could potentially cause gravitational settling and sequestering of the water. Two, dust may be sufficiently radiatively active to prohibit clouds from forming and potentially allowing a greater degree of water vapor transport out of the polar region. The MGS TES instrument also spans the wavelength region over which Martian dust is absorbing, allowing its retrieval.

Surface Temperature. The north polar region surface temperature during the northern polar season can be compared to the Viking era to further elucidate potential interannual differences and to understand the vapor, water-ice, and dust retrievals. MGS TES is also able to accomplish this, typically via the 30-micon radiance. Surface temperatures can confirm that carbon dioxide is not present on the cap surface, implying that the brightening is likely water ice.

Intercorrelation of data: To date, there has been no comprehensive study to understand the partitioning of water into vapor and ice clouds, and the associated effects of dust and surface temperature in the north polar region. Ascertaining the degree to which water is transported out of the cap region versus within the cap region will give much needed insight into the overall story of water cycling on a seasonal basis. In particular, understanding the mechanism for the polar cap surface albedo changes would go along way in comprehending the sources and sinks of water in the northern polar region. We approach this problem by examining TES atmospheric and surface data acquired in the northern summer season and comparing it to Viking data when possible. Because the TES instrument spans the absorption bands of water vapor, water ice, and dust, and has a clear window to the surface near 30 microns, surface temperature and all three aerosols can be retrieved simultaneously. This presentation will be a status update on our current analysis and interpreta-

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